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## On the Calculation of the Seismic Parameter $\phi$ at High Pressure and High Temperatures

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Comparison of the Murnaghan equation of state with the Birch equation shows that, for a given value of pressure, the values of  $(\rho/\rho_0)$  calculated from the two equations differ by less than 1% to a pressure equal to 0.5  $K_0$  (where  $K_0$  is the zero-pressure isothermal bulk modulus), but the corresponding values of the seismic parameter  $\phi$  differ by 10%. The value of  $\phi$  is extremely sensitive to the choice of the equation of state because  $\phi$  is the derivative of pressure with respect to density. The good agreement between the two equations of state for pressure as a function of density observed for some materials does not imply the same agreement in the relationship between  $\phi$  and pressure. Expressions for  $\phi(P)$  that take into account the first order nonlinear dependence of the bulk modulus on pressure are presented, and their applications are discussed. Temperature correction of the pressure-dependent  $\phi$  is also considered.

Comparison of the seismic parameter  $\phi_{\text{LAB}}$ , determined in the laboratory for various materials, with the values actually observed in the field,  $\phi_{\text{FLD}}$ , can be used to estimate the composition of a homogeneous isothermal layer within the earth. If a particular equation of state is assumed, then the seismic parameter may be written as a function of pressure because the definition of the adiabatic bulk modulus  $K_s$  implies that

$$\phi = \left(\frac{\partial P}{\partial \rho}\right)_{s} \tag{1}$$

Birch [1939] used the Murnaghan theory of finite strain to calculate the rate of change of seismic velocities with pressure. O. L. Anderson presented an equation for a pressure-dependent  $\phi$  based on the Murnaghan equation of state and illustrated its applicability at high pressure. He concluded [O. L. Anderson, 1966, p. 730] that 'Birch's equation of state, in its form which is appropriate to a general value of  $K_0$ , leads to essentially the same results as does the Murnaghan equation'; we believe the two equations lead to different results.

In this paper we compare the values calculated for  $\phi$  from both the Murnaghan and the Birch equations of state and discuss the sensi-

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tivity of  $\phi(P)$  to the choice of the equation of state; we believe the Birch form superior to that of Murnaghan. Expressions for  $\phi(P)$  that take into account the first-order nonlinear dependence of the bulk modulus on pressure are given, and their implications are discussed. Correction of the pressure-dependent  $\phi$  for temperature is considered.

## SENSITIVITY OF $\phi(P)$ TO THE CHOICE OF EQUATION OF STATE

The equations of state most widely used in geophysics are those of Murnaghan [1944, 1949] and Birch [1939, 1947, 1952]. We examine the dependence of  $\phi(P)$  on the form of the equation of state used to describe the elastic behavior of solids.

The Murnaghan equation of state is derived from the assumption that bulk modulus is a linear function of pressure:  $K(P) = K_0 + mP$ , where  $K_0$  is the adiabatic bulk modulus evaluated at zero pressure, and m is a material constant defined by  $m = \{(\partial K/\partial P)_*\}_{P=0}$ . Since  $K = \rho(dP/d\rho)$ ,

$$P_M = (K_0/m)[(\rho/\rho_0)^m - 1]$$
 (2)

The subscript M denotes parameters calculated from the Murnaghan equation of state.

The Birch equation of state, derived from Murnaghan's theory of finite strain [Murnag-

han, 1951] with cubic and quadratic terms of strain retained in the Helmholtz free energy [Birch, 1947, 1952], leads to

$$P_B = (3K_0/2)[(\rho/\rho_0)^{7/3} - (\rho/\rho_0)^{8/3}]$$

$$\cdot \{1 + (\frac{3}{4})(m-4)[(\rho/\rho_0)^{2/3} - 1]\}$$
 (3)

The subscript B refers to the parameters calculated from the Birch equation of state.

From equation 2, we find the derivative

$$dP_M/d\rho = \phi_0(\rho/\rho_0)^{m-1} \tag{4}$$

where  $\phi_0 = (K_0/\rho_0)$ . To express  $\phi$  as a function of pressure, we substitute equation 2 in the form

$$\rho/\rho_0 = [1 + m(P_M/K_0)]^{1/m}$$
 (5)

and obtain

$$\mathcal{O}_{M} = dP_{M}/d\rho 
= (K_{0}/\rho_{0})[1 + m(P_{M}/K_{0})]^{(m-1)/m}$$
(6)

Equation 6 corresponds to equation 8 of O. L. Anderson's [1966] paper, and it is noted that he derived this expression in a different way.

Similarly, from the Birch equation of state, we have

$$P_B = (3K_0/2)y^5\{(y^2-1) + b_1(y^2-1)^2\}$$
 (7)

nd

$$\phi_B = dP_B/d\rho = (\phi_0/3)\{3y^4[1 + 2b_1(y^2 - 1)] + (5/y^3)(P_B/K_0)\}$$
(8)

where  $y = (\rho/\rho_0)^{1/3}$  and  $b_1 = (3/4)(m-4)$ . To obtain  $\phi_B$  as a function of pressure, the Birch equation of state must be solved numerically for  $(\rho/\rho_0)$  as a function of pressure.

It has previously been recognized that the Murnaghan equation 2 will be limited to values of  $P < 0.5K_0$  in estimating  $(V/V_0)$  [e.g., O.L. Anderson, 1968, p. 170]. We show below that its validity for  $\phi$  does not extend as high as  $P \simeq 0.5K_0$ .

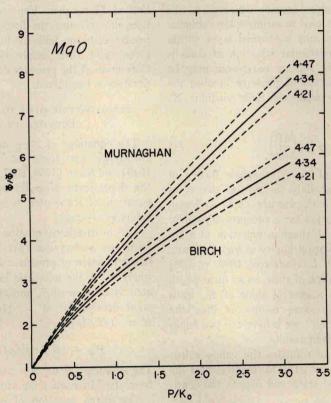


Fig. 1. Comparison of the seismic  $\phi$  calculated from the Birch and the Murnaghan equations for periclase (at 298°K).